



Radar-derived bed roughness characterization of Institute and Möller ice streams, West Antarctica, and comparison with Siple Coast ice streams

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[1] Subglacial bed conditions exert a significant control on ice stream behavior and evolution, and can be characterized by determining bed roughness from FFT analysis of radar-imaged basal reflectors. Here we assess bed roughness across Institute and Möller ice streams, West Antarctica, and compare our findings with bed roughness determined across the Siple Coast ice streams. We find that variations in bed roughness are spatially organized, and attribute this to the varying efficacy of subglacial erosion and deposition, with rougher (inland, slow-flowing) regions largely manifesting preglacial topography, and smoother (downstream, fast-flowing) regions evincing significant postglacial modification to the subglacial landscape. The observed similarities between bed roughness characteristics of IIS/MIS and the Siple ice streams suggest that IIS and MIS are largely underlain by wet, poorly consolidated sediments, and may therefore be vulnerable to the types of dynamical instabilities experienced by the Siple ice streams. **Citation:** Bingham, R. G., and M. J. Siegert (2007), Radar-derived bed roughness characterization of Institute and Möller ice streams, West Antarctica, and comparison with Siple Coast ice streams, *Geophys. Res. Lett.*, 34, L21504, doi:10.1029/2007GL031483.

1. Introduction

[2] Elucidating the nature of the subglacial interface beneath an ice sheet is critical to determining controls on ice dynamics and the configuration and stability of regions of streaming flow. This is especially pertinent for the West Antarctic Ice Sheet (WAIS), where some ice streams evince both spatial migration and dramatic temporal changes in discharge [Alley and Bindschadler, 2001; Conway *et al.*, 2002]. Much of our knowledge of basal conditions beneath the WAIS is derived from studies conducted over the Siple Coast region (Figure 1), where ice streams separated by interstream ‘ridges’ of slow-flowing ice feed the Ross Ice Shelf. Airborne radio-echo sounding (hereafter RES), seismic analyses, and borehole sampling conducted across this region have revealed that ice streams are underlain extensively by unconsolidated marine sediments, emplaced when the ice sheet was smaller than today. These sediments

deform readily when wet facilitating streaming ice flow above despite low driving stresses [Anandakrishnan *et al.*, 1998; Bell *et al.*, 1998; Tulaczyk *et al.*, 2000; Peters *et al.*, 2006].

[3] By contrast, little is known about the flow and basal characteristics of the ice streams that feed the Filchner-Ronne Ice Shelf (FRIS; locations in Figure 1). Institute Ice Stream (IIS; 81.5°S, 75°W) drains one of the largest catchments in the WAIS, disgorging 22.7 ± 2 Gt a⁻¹ of mass to the FRIS [Scambos *et al.*, 2004], yet its flow is poorly constrained due to a combination of remoteness and sparse satellite velocity measurements. The existing studies that have discussed ice flow through the region [Drewry *et al.*, 1980; Scambos *et al.*, 2004; Joughin *et al.*, 2006] suggest it may be underlain by weak till, comparable to the Siple ice streams, and therefore the inherent instabilities present in those ice streams might also apply to IIS and flanking regions draining to the FRIS.

[4] Bed roughness, the vertical variation of the subglacial interface with horizontal distance [Taylor *et al.*, 2004], provides a useful first-order summary measure of regional basal characteristics, yielding critical regional-scale insights into ice stream behavior and controls [Bingham *et al.*, 2007; Siegert *et al.*, 2004, 2005]. Here, we calculate bed roughness across IIS and the neighboring Möller Ice Stream (MIS, 82.3S, 65°W) using RES-captured images of the basal reflector. We compare our findings with those from the better known Siple Coast, and use the results to assess the flow, basal characteristics and stability of IIS and MIS, key contributors of the WAIS to the FRIS.

2. Methods

[5] In recent years several approaches have been adopted to characterize bed roughness across parts of Antarctica. One approach is to analyze the detailed surface topographic data now available from satellite altimetry, using transfer functions to invert for variations in basal topography [Thorsteinsson *et al.*, 2003; Joughin *et al.*, 2006]. This method is most effective in regions whose surface velocity fields are well characterized by remote sensing, but 69% of the combined IIS/MIS catchment lies outside current satellite-derived surface velocity fields (Figure 1). Alternative methods utilize bed topography directly imaged by RES. One such approach is to interpolate a two-dimensional basal topographic surface between raw bed returns collected along RES tracks, and calculate the mean and standard deviation of the elevation of each grid cell. Variations in elevation between grid cells can thereon be used to define bed roughness [Rippin *et al.*, 2006]. This method is only viable where RES tracks follow a closely-spaced grid network, which is

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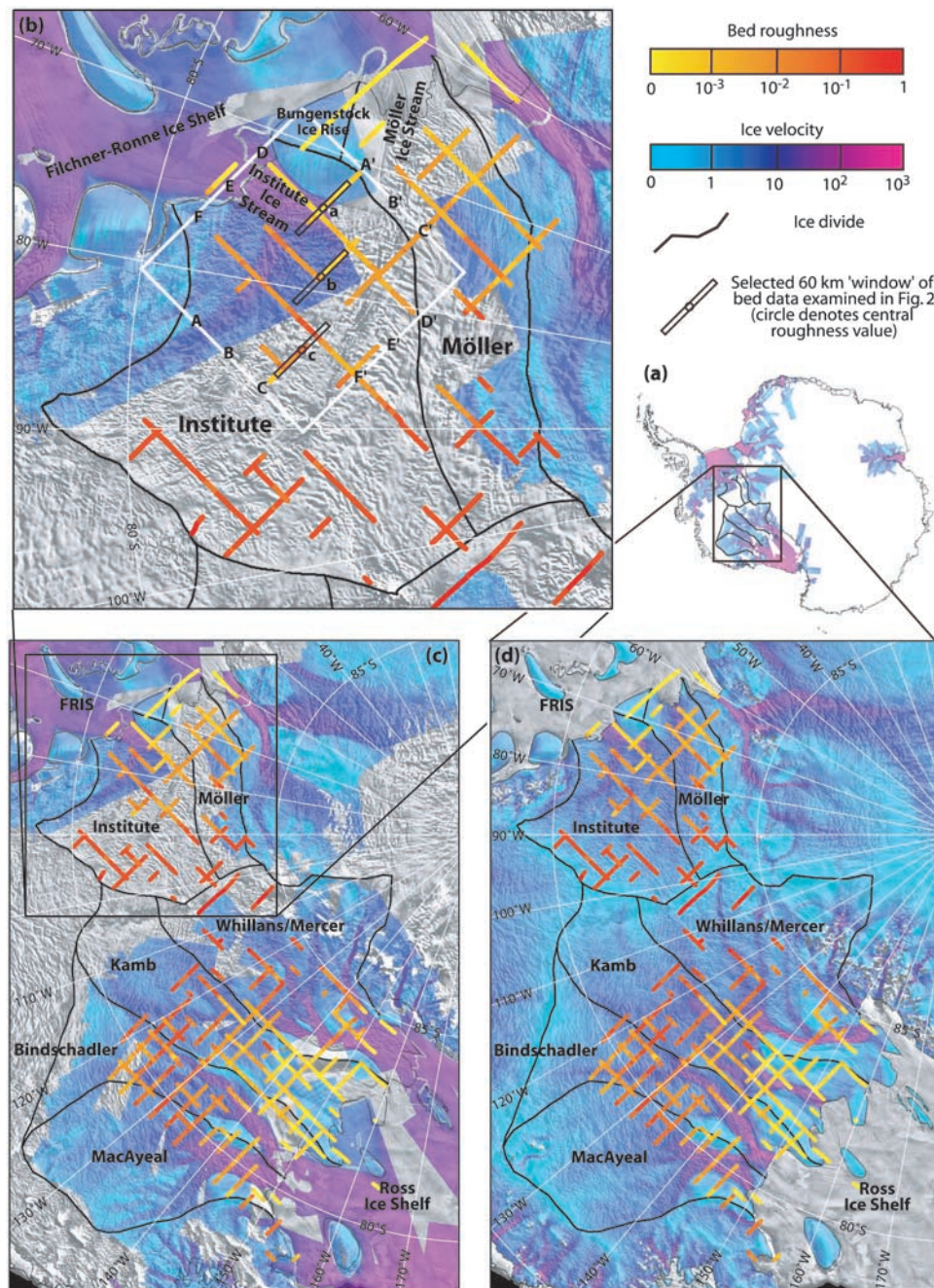


Figure 1. Charts of total bed roughness along RES lines for catchments draining to the Filchner-Ronne (FRIS) and Ross Ice Shelves. (a) Catchment locations superimposed over surface ice velocities [after *Joughin et al., 2006; Rignot, 2006*]. Along-RES-track bed roughness superimposed over MOA imagery [*Haran et al., 2006*], over which is draped either the (b) and (c) available coverage of satellite-derived surface ice velocities or (d) depth-averaged balance velocities [after *Le Brocq et al., 2006*]. In Figure 1b, a, b, and c mark selected elevation profiles analyzed for bed roughness power spectra in Figure S1. The white boxed area and upper-case letters locate bed roughness profiles analyzed in Figure S2.

not the case for the available data over IIS and MIS (Figure 1). The remaining alternative is to interrogate RES data directly along-track to extract the spectral roughness of subglacial topography [*Taylor et al., 2004; Siegert et al., 2004, 2005; Bingham et al., 2007*]. To date, limited satellite coverage and the remoteness of the IIS/MIS region have deferred an analysis of its basal properties, yet bed returns are available from a wide-ranging airborne RES survey con-

ducted across the region between 1977 and 1979. Here we calculate bed roughness directly along these RES flight tracks spanning IIS and MIS.

[6] The RES data analyzed here were obtained by the Scott Polar Research Institute/National Science Foundation/Technical University of Denmark (SPRI-NSF-TUD) consortium between 1977 and 1979. Details of the surveys, including navigation and errors, are provided in the auxil-

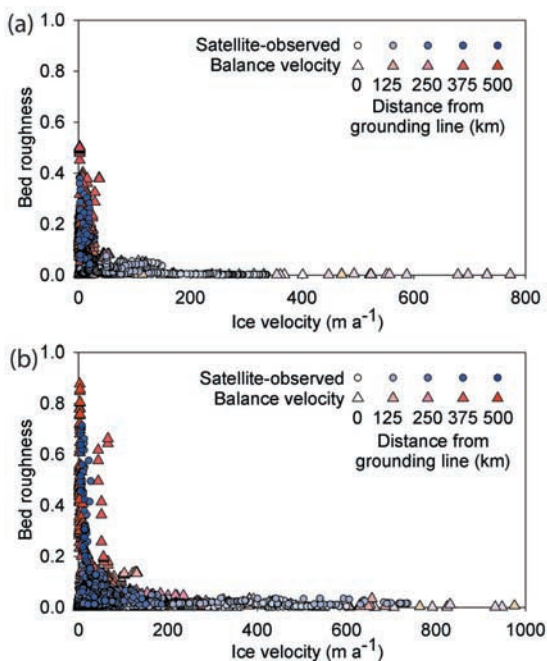


Figure 2. Dimensionless bed roughness values calculated along RES lines in Figure 1 compared with satellite-observed ice velocities (as mapped in Figure 1c) and depth-averaged balance velocities (as mapped in Figure 1d) across the (a) Institute and Möller and (b) Siple Coast catchments (Whillans/Mercer, Kamb, Bindschadler, and MacAyeal as delineated in Figures 1c and 1d). The symbols are shaded such that darker symbols are close to the grounding line and lighter symbols are far from the grounding line.

ary materials [see also *Drewry*, 1983; *Siegert*, 1999].¹ Critically, they comprise the only RES data ever collected over the remote IIS and MIS, and as such provide the only data that can be used directly to characterize their basal properties. The quality of the SPRI-NSF-TUD data has been compared favorably elsewhere in Antarctica with modern datasets [*Siegert et al.*, 2004; *Bingham et al.*, 2007]. Over 7000 km of flight track are available over the IIS/MIS region; the bed is imaged in 72% of 4117 shots analyzed.

[7] The roughness of the basal reflector was calculated using a Fast Fourier-Transform- (FFT-) based technique, which characterizes the wavelength-related vertical variation of a surface with horizontal distance [*Taylor et al.*, 2004]. Each FFT was applied over a moving 60-km window, bed roughness being the integral of the resulting FFT spectral power density plot for each analysis window (for examples and further information, see auxiliary materials, section S2). Each bed roughness value is plotted at the centrepoint of its analysis window, providing a spatial map of bed roughness variations along flight tracks over the WAIS (Figure 1). We compare the roughness variations with satellite-derived ice-surface velocities [after *Joughin et al.*, 2006] and depth-averaged balance velocities [after *Le Brocq et al.*, 2006]. We consider both of these velocity datasets because the satellite-derived data cover only 31% of the IIS/

MIS region (Figure 1b), thus balance velocities can be used as proxies for ice flow elsewhere.

3. Results

[8] Across the Institute and Möller catchments total bed roughness ranges between 0.0002 and 0.504, averaging 0.057, comparable to values measured across the Siple Coast [*Siegert et al.*, 2004]. The roughness values exhibit clear spatial organization (Figure 1). Beneath the main trunk of IIS the basal reflector is remarkably smooth, never rising above 0.01 between points D and a on Figure 1b. Beneath the downstream portion of MIS too, the basal reflector is relatively smooth, remaining below 0.03 downstream of point C' (Figure 1b). Away from the regions of active streaming, and with distance from the grounding line towards the ice divides, the basal reflector becomes progressively rougher. These findings are true regardless of whether the roughness windows analyzed are quasi-parallel or quasi-orthogonal to ice flow, and mirror spatial variations in bed roughness found across the Siple Coast (see Figures 1c and 1d). Further support for this statement is given in auxiliary materials, section S3.

[9] In Figure 2, bed roughness is plotted versus ice velocity. Only very low bed roughness values are attained where ice speeds are high, whereas high bed roughness is only measured where ice speeds are low (Figure 2a). Typically, away from the grounding line, where ice flows more slowly, bed roughness is high. Towards the grounding line, where ice velocities are more variable but can be very high, bed roughness is low (Figure 2a).

[10] Plotting roughness values directly against distance from the grounding line (Figure 3) reveals the ice stream beds become progressively smoother towards the grounding line, such that $y = ae^{bx}$, where y = bed roughness (dimensionless), a and b are constants that vary between individual catchments, and x = distance from the grounding line (km; derivation outlined in auxiliary materials, section S4). For both IIS and MIS the coefficient of determination (R^2) for this relation is >0.5 , thus $>50\%$ of the variation in roughness is 'explained' by the distance of the roughness datapoint from the grounding line. Conducting the same exercise for the four Siple catchments delineated in Figure 1c confirms that the same relation also applies strongly over the Whillans/Mercer and Kamb catchments (Figures 3c and 3d), and is especially strong for the Kamb catchment, yielding an R^2 of 0.77. The same relationship is weaker over the Bindschadler and MacAyeal catchments ($R^2 < 0.5$), probably in the latter case due to insufficient data coverage in upstream parts of the catchment (Figures 3e and 3f).

4. Discussion

[11] Our results demonstrate bed roughness varies in a distinct spatially-organised manner beneath the two largest catchments feeding the FRIS. The roughness trends provide unique insights into the subglacial environment across this poorly known region. The parallel with the general pattern of roughness variations observed across the Siple Coast region is striking, and suggests strongly that the subglacial environments beneath the two regions are analogous. Beneath the Siple Coast ice streams smooth basal reflectors

¹Auxiliary material data sets are available at <ftp://ftp.agu.org/apend/gl/2007gl031483>. Other auxiliary material files are in the HTML.

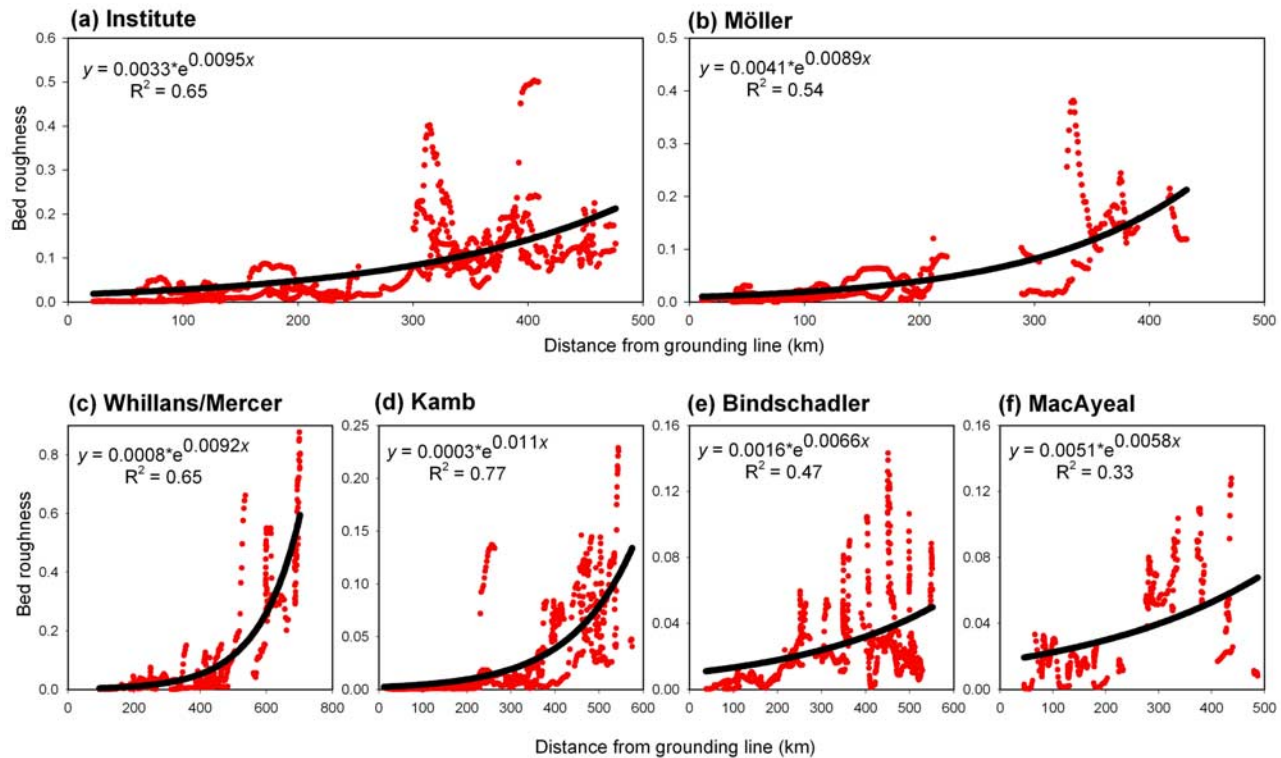


Figure 3. Plots of bed roughness versus distance from the grounding line for (a)–(b) the Institute and Möller catchments and (c)–(f) the Siple catchments.

have been linked to the subglacial existence of unconsolidated sediments that deform readily when wet, allowing fast flow to persist despite low driving stresses [Engelhardt and Kamb, 1998; Tulaczyk et al., 2000; Blankenship et al., 2001; Peters et al., 2006]. IIS also exhibits fast flow and low driving stresses [Scambos et al., 2004; Joughin et al., 2006] together with a smooth basal reflector (Figure 1a). Hence the most likely explanation for this combination of phenomena is a widespread unconsolidated sediment layer beneath IIS and much of its catchment. There is no increase in bed roughness moving eastwards from the lower 200 km of IIS across the slow-flowing Bungenstock Ice Rise and into the downstream 200 km of MIS (Figure 1a; see also plot AA' in Figure S2), and this likely reflects a thick basal sediment layer pervading most of the IIS/MIS region fringing the FRIS.

[12] Subglacial roughness properties beneath the IIS and MIS catchments are most analogous to those beneath the Kamb and Whillans/Mercer catchments. All four catchments are characterised by a distinct roughening of subglacial topography with distance inland (Figure 3) and with downstream beds uniformly smooth regardless of whether they underlie contemporary ice streams or interstream ridges. The similarity is significant, because the Kamb Ice Stream is demonstrably prone to unstable behaviour, presenting evidence of several earlier periods of fast flow and lateral migration preceding its current stagnant state [Retzlaff and Bentley, 1993; Bentley et al., 1998; Kamb, 2001; Catania et al., 2003]. A recent slowdown of Whillans Ice Stream suggests it too is prone to similar dynamic behaviour [Bougamont et al., 2003]. Key to the inherent instability of the Kamb and Whillans ice streams is thought

to be a thick underlying deformable sediment layer and an absence of constraining subglacial topography [Kamb, 2001; Bougamont et al., 2003; Siegert et al., 2004]. The ubiquity of the sediments tends to dampen subglacial topographic protuberances, leaving few constraints to lateral migrations of the ice streams; much of the unstable behaviour has instead been linked to fluctuating thermal conditions in the sediments which deform readily when wet but stiffen when frozen [Tulaczyk et al., 2000]. The inference that a thick swathe of deformable basal sediments also underlies IIS and MIS suggests therefore that these two FRIS ice streams are likely vulnerable to similar dynamic instabilities.

[13] Boreholing to the beds of Whillans, Kamb and Bindschadler ice streams and their interstream ridges has revealed the subglacial tills beneath the region are primarily composed of Tertiary marine sediments [Engelhardt et al., 1990; Engelhardt and Kamb, 1998; Kamb, 2001]. Aero-geophysical surveying has been used to assess the likely distribution of the sediments over the Siple Coast region by isostatically adjusting the subglacial topography to demarcate the paleo-shoreline, below which marine sedimentation would have been possible prior to the formation of the WAIS and/or during periods of ice-sheet collapse [Studinger et al., 2001]. The areas with the most sediment correspond with smooth basal reflectors, suggesting 'smoothness' can be used elsewhere as an indicator of subglacial marine sediments. A physical explanation is provided by considering rougher beds as legacies of a mountainous preglacial landscape, and smoother beds as symptomatic of its progressive burial beneath marine sediments. We therefore interpret the widespread region of smooth subglacial topography underlying the lower sectors of IIS and MIS as an

extensive covering of marine sediments emplaced when the ice-sheet was less advanced than today.

[14] The existence of such a smooth deforming basal sediment layer would offer no basal constraint to previous or future lateral migrations of IIS and MIS, and does not discount the possibility that they may be subject to periods of stagnation. The possibility of such unstable behavior has implications for the history and future stability of the FRIS, to which these ice streams, especially IIS, are significant contributors. We therefore support the contention of Scambos *et al.* [2004] that sampling of basal sediments beneath IIS would offer valuable scientific insights. If the basal material is marine, and was therefore emplaced when the ice sheet was smaller than today, it could be dated to constrain when the region was last glaciated, providing critical insights into the age and stability of the FRIS and the WAIS.

5. Conclusions

[15] Applying FFT analysis to airborne radar data collected over Institute Ice Stream (IIS) and Möller Ice Stream (MIS), West Antarctica, we have determined spatial variations in the roughness of the bed that correspond with variations in ice flow and distance from the grounding line. Typically, the bed is smoother where ice flows faster and towards the grounding line, whereas the bed is rougher where ice flows more slowly and towards the ice divides. The variations in bed roughness most likely reflect differences in the efficacy of erosion and deposition beneath the ice. They mirror similar variations beneath the Siple Coast ice streams, and are most similar to bed roughness characteristics across the Kamb and Whillans/Mercer catchments. Across the downstream sectors of IIS and MIS and the intervening Bungenstock Ice Rise the basal reflector is contiguously smooth. By analogy with earlier Siple Coast-based studies, we contend that IIS and MIS are widely underlain by unconsolidated marine sediments, and therefore these key contributors to the Filchner-Ronne Ice Shelf may be vulnerable to the types of dynamical instabilities experienced by the Siple Coast ice streams.

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