

1 **Circum-Arctic Lithosphere-Basin Evolution: An Overview**

2 **Larry S. Lane and Randell A. Stephenson**

3
4 **Abstract**

5 A new collection of papers spanning the breadth of the Arctic provides new insight into the
6 region’s geodynamic evolution. New results pertain to lithospheric structure, the link between
7 magmatic and extension-related tectonic processes, variations in the composition and velocity
8 structure of the lower crust in the Amerasia Basin, and provenance and paleogeography of
9 Paleozoic to Triassic successions across the Arctic. Elucidation of geodynamic processes in the
10 Eurasia Basin suggests new hypotheses for future research in the complex and poorly understood
11 Amerasia Basin. New results from detrital zircon provenance studies as well as from
12 stratigraphic facies compilations constrain the Late Paleozoic to Triassic paleogeography of the
13 Arctic realm.

14
15 **Key Words**

16 Circum-Arctic; tectonics; geodynamics; paleogeography; magmatism

17
18 **1. Introduction**

19 This circum-Arctic lithosphere-basin Special Issue evolved from a Special Session at the
20 Geological Society of America Annual meeting in Vancouver, Canada in 2014. The session’s
21 aim was to highlight new research across the Arctic, reflecting the breadth of lithospheric

22 interactions that occurred over the past ~700 Myr, emphasizing thematic linkages, such as inter-
23 regional, onshore-offshore, lithosphere-basin, magmatism-tectonism and so on. The modern
24 Arctic is the product of a long history of plate tectonic processes that closed and opened
25 Phanerozoic oceans, deformed sedimentary basins and created orogens, together with surface
26 processes that produced and transported sediments from orogenic highlands into adjacent
27 sedimentary basins. By assembling a multidisciplinary suite of papers into one Special Issue, this
28 project attempts to facilitate greater understanding of the interdependence of the processes
29 underlying the tectonic, magmatic and sedimentary evolution of the Arctic, to encourage new
30 avenues for future multidisciplinary studies, and possibly to foster fresh insights into the linkages
31 between these processes.

32

33 From the 1960s and until the 1990s, geoscientific research in the Arctic was driven largely by
34 cyclical activity in the petroleum exploration sector, interspersed with intermittent public sector
35 and academic programs. The Arctic's remoteness and inaccessibility was, and continues to be, an
36 impediment to research. Although these efforts resulted in many important advances in
37 understanding the region's geological history, it remained poorly understood. Over the past 10-
38 15 years, the evolution of the Arctic has been increasingly a focus of geoscience research.

39 Initially, this was driven by renewed exploration investments by the petroleum sector. However,
40 as technological advances in petroleum production reduced the economic incentive for Arctic
41 resource development, a new incentive for research arose in the United Nations' 1982
42 Convention on the Law of the Sea (UNCLOS), Article 76, which sets out how coastal states
43 define the outer limits of their legally defined continental shelf beyond 200 NM, where one
44 exists (e.g., United Nations, 1999). The criteria set out in Article 76 required coastal states to

45 collect significant volumes of bathymetric and sediment thickness data across the Arctic
46 continental margins, which provided an opportunity to acquire new crustal structure data. It was
47 also a challenge to drive new technologies (e.g., new seismic sources designed to operate in
48 heavy ice conditions) , data collection methods (e.g., multi-national, multi-ship ice-breaker
49 operations) and international collaborations in order to be successful. Data for several of the
50 contributions to this Special Issue were acquired through surveys in support of programs to
51 define nations' extended continental shelves, as required under Article 76 of the Convention.

52

53 Meanwhile, onshore, a dramatic increase in the volume of research based on U-Pb age analyses
54 of detrital zircon began to provide new insights into the provenance histories of multiple
55 Paleozoic and Mesozoic depositional systems around the Arctic. In many cases, they identified
56 previously unrealized common sources for sands deposited into sedimentary basins that were
57 subsequently tectonically dismembered and dispersed. In a few cases, age populations occur for
58 which no plausible sources are known. Further, as the technique has become more widespread, it
59 is increasingly apparent that sedimentary recycling is more common than previously understood.
60 Thus, the immediate zircon sources may be quite different from the original magmatic sources.
61 Such cycles of inheritance can be traced back through several generations. In combination with
62 stratigraphic and structural analyses, the use of detrital zircon geochronology continues to
63 advance our understanding of Arctic paleogeography across the region and throughout the
64 Phanerozoic. Three contributions in this issue are based on detrital zircon geochronology.

65

66 **2.0 Thematic Overview**

67 Altogether, the contributions in this Special Issue reflect current research in the Arctic realm
68 encompassing multiple aspects of the geosciences: lithospheric structure, magmatic processes,
69 rheology, tectonic evolution, and provenance and paleogeography. Geographically, the 15
70 presented studies span much of the Arctic region (Fig. 1). In Figure 1, they are colour coded
71 according to two broad categories: lithospheric structure and geology. The former is largely
72 marine and remotely sensed, further differentiated in Figure 1 based on their principal data sets
73 (seismology vs potential field). The latter consists of land-based field studies. Here we
74 summarize the contributions in the Special Issue. In order to facilitate the presentation, we have
75 organized what follows into these two broad themes: Lithospheric structure and Geology (Fig.
76 1).

77

78 *2.1. Lithospheric Structure*

79 The first theme, lithospheric structure, is explored primarily using seismology supplemented
80 with potential field modelling. Despite the logistical complexities of acquiring seismic data in
81 heavy ice, these new studies determine the physical characteristics of its crust with sufficient
82 resolution to make new inferences about tectonic setting and internal structure.

83

84 In an important and comprehensive study of the crustal structure in southern Canada Basin,
85 Chian et al. (2016) employ a suite of sonobuoy records supported by multichannel seismic
86 reflection profiles and potential field modelling to classify crustal types within the basin. They
87 distinguish three crustal types mainly on the basis of their P-wave velocity analysis: thinned
88 continental crust (5.5-6.6 km/s), oceanic crust (6.7-7.1 km/s), and high-velocity transitional crust

89 (7.2-7.6 km/s). A key finding is that oceanic crust occupies a broadly pentagonal area within
90 central Canada Basin some 600 km long by 350 km wide. This publication undoubtedly will be
91 used widely to underpin future tectonic syntheses of the basin.

92

93 In a new 400 km seismic refraction and gravity transect from the Greenland margin of Nares
94 Strait southward into northern Baffin Bay, Altenbernd et al. (2016) present a new crustal
95 structure model, imaging the continent-ocean transition in this area. Crustal layers are defined in
96 the continental, transitional and oceanic portions of the transect. Velocities of 7.7 km/s
97 characterize the sub-oceanic mantle; such low velocities are inferred to relate to serpentinization.
98 Serpentinization as well as thin oceanic crust in this area are correlated with slow spreading
99 rates. Moho depths vary from 34 km beneath Greenland to 13 km beneath the oceanic crust.
100 No evidence was found to suggest significant magmatism occurred within the continent-ocean
101 transition zone.

102

103 Anudu et al. (2016) applied a series of derivative filters to regional aeromagnetic and gravity
104 anomaly datasets, to enhance signals of magmatic and structural features across the Canadian
105 Polar margin adjacent to Alpha Ridge. They recognize five magnetic domains based on anomaly
106 patterns and attributes. They infer that Cretaceous to Paleogene magmatic bodies are more
107 prevalent than the existing mapping suggests; that the marginal basin underlying the continental
108 shelf post-dates Cretaceous magmatism, and that most of the intrusions were emplaced in a NW-
109 SE minimum compressive stress field, believed to correlate with Early Cretaceous High Arctic
110 Large Igneous Province (HALIP) magmatism.

111

112 Oakey and Saltus (2016) present magnetic and gravity models along regional transects through
113 the Alpha-Mendeleev Ridge complex, which extends over a region exceeding one million square
114 kilometres, and compare the resulting crustal characteristics with other major large igneous
115 provinces. They conclude that HALIP magmatism was largely or completely intruded into
116 continental crust that was intensively intruded and underplated by mafic magmatism. They refer
117 to the continental Kerguelen Plateau as an appropriate analogue.

118

119 Using seismology and potential field analyses, several studies focus on the magmatic processes
120 associated with ocean basin tectonic evolution in the context of mid-ocean ridge processes and
121 large igneous provinces (LIPs). These contributions provide detailed seismic and potential field
122 interpretations along the Greenland continental margins and adjacent oceanic spreading ridges
123 that elucidate the processes involved in magmatic and amagmatic formation of oceanic crust in
124 the Eurasian Basin and north Atlantic during Cenozoic time.

125

126 A contribution by Schmidt-Aursch and Jokat (2016) uses 3D gravity modelling in parts of the
127 Eurasian Basin to resolve variations in crustal thickness related to temporal and spatial variations
128 in magma supply to the ultra-slow spreading Gakkel Ridge. They describe three regional
129 domains: two that are magma-rich, separated by a central domain containing areas of sparse but
130 focused volcanism. Magnetic anomaly amplitudes vary in complex ways, both along and across
131 the study area. However, they appear to be uncorrelated with crustal thickness variations. In the

132 western domain a period of more robust magmatism with well-defined anomalies is inferred to
133 correlate with northward propagation of a chemically distinct North Atlantic mantle.

134

135 In a related study, interpretation of aeromagnetic surveys in the north Greenland, Fram Strait and
136 Lena Trough areas, Jokat et al. (2015) have clarified multiple relationships between magmatic
137 events and tectonic evolution adjacent to northern Greenland. First, the latest Cretaceous Kap
138 Washington volcanics, extensively exposed onshore, are more widely preserved offshore than
139 was previously recognized. They also recognize that the Arctic spreading centre required ~11
140 Myr to propagate southward from Fram Strait to the Spitzbergen Fracture zone. Also, the ultra-
141 slow spreading centre in Lena Trough, inferred to be non-volcanic, is characterized by a subdued
142 magnetic signature that may reflect magnetization of exhumed serpentized mantle.

143

144 A new crustal seismic refraction transect and coincident gravity profile, of some 500 km length,
145 extends from Kong Oskar Fjord into the Atlantic Ocean off east Greenland (Hermann and Jokat,
146 2016). It presents a new model of the continent-ocean transition in this area, characterized by a 3
147 km thick lower-crustal high-velocity zone that is inferred to indicate excess magmatism in the
148 vicinity of the Jan Mayen Fracture Zone, dating to the separation of the Jan Mayen
149 microcontinent from Greenland in Oligocene time. This implies that fracture zones may play an
150 important role in the production and distribution of magma at extended continental margins.

151

152 The response of the lithosphere to applied tectonic stresses is dependent on its rheology. Local
153 and regional variations determine if and where strain becomes concentrated, and also the
154 mechanism(s) by which the strain occurs. External influences on rheology are the thermal
155 environment, the lithology and the strain-rate.

156

157 In the Barents and Kara Sea region, Gac et al. (2016) employ a 3D lithospheric-scale structural
158 and thermal model to compute an effective elastic thickness map and rheological model of the
159 region. They illustrate the correlation between lithospheric thickness, lithospheric strength and
160 effective elastic thickness. Also, the region has an asymmetric lithospheric structure, thinner and
161 hotter in the west, colder and thicker in the east. Thus differing responses to imposed tectonic
162 stresses are to be expected across the region. Computed gravitational potential energy across the
163 region indicates that ridge-push is may be a sufficient mechanism to produce contraction on
164 western Barents shelf.

165

166 Li et al. (2016) used backstripping and 2D subsidence modelling to reproduce the thermal
167 subsidence history of the southern margin of Canada Basin in comparison to an analogous area
168 within the South China Sea. Both settings involve hyperextended crust and deep basins with
169 crustal geometry constrained by seismic reflection profiles and gravity data. The authors
170 conclude that a weak lower crust and high heat flow led to a lower crust that extended much
171 more than the overlying crust, implying substantial ductile thinning within the lower crust.

172

173 2.2. *Geology*

174 Studies of stratigraphic and structural history, integrated with geochronology and geochemistry
175 of magmatic rocks provides additional information on the deeper lithospheric processes during
176 tectonism, as well as on regional tectonic evolution, through time.

177

178 Within the context of the temporally and spatially complex mid-Paleozoic tectonic history of the
179 Timan Range of Arctic Russia, Pease et al. (2016) attempt to resolve ambiguities in the tectonic
180 setting of the magmatism by examining the petrology and geochemistry of Devonian basalts.
181 Chemical (including trace and rare-earth elements) and isotopic relationships indicate that the
182 low-potassium tholeiitic magmas comprise two discrete groups, derived from mid-ocean ridge
183 basalts, distinguished by differing amounts of contamination by continental crust. The
184 magmatism is interpreted to result from local melts related to rifting of a mainly non-volcanic
185 continental margin sampling a heterogeneous mantle source. It is distinct from back-arc
186 extension or mantle plume-related settings and suggests a complex interplay between tectonic
187 settings within and marginal to the East European Craton.

188

189 In an ambitious review and synthesis paper, Moore and Box (2016) distill the Phanerozoic
190 tectonic evolution of Alaska into a suite of 14 time-slice maps summarizing the evolution of
191 nearly two dozen discrete tectonostratigraphic terranes, in three regional domains. Alaska's
192 location at the intersection of the Pacific and Arctic realms of the Laurentian craton has
193 positioned it to interact with exotic terranes derived from both directions throughout Phanerozoic
194 time. Accordingly, its stratigraphic, magmatic and tectonic evolution is complex and diverse. It

195 incorporates terranes with affinities to the Siberian, Baltican and Laurentian cratons, as well as to
196 various oceanic realms. Moore and Box summarizes the current state of knowledge for the bulk
197 of Alaska and, with its extensive citation list, provides an entry point into the Alaskan geoscience
198 literature for readers.

199

200 Studies of provenance and paleogeography provide information that permits correlations within
201 continents and between continents and terranes. Such linkages may help to characterize terranes
202 that are otherwise poorly understood. Four papers in this issue focus on these topics, of which
203 three are based on detrital zircon (DZ) studies. One study synthesizes regional sedimentology
204 and stratigraphy, supported by previously published DZ results.

205

206 Zhang et al. (2015) report on the provenance and paleogeography of a latest Mesoproterozoic
207 and Neoproterozoic metasedimentary succession located on the northern margin of the Baltic
208 Shield, now imbricated within Caledonian nappe complexes. Using new DZ data combined with
209 facies analyses and an extensive dataset of paleocurrent data, they reconstruct the
210 Neoproterozoic paleogeography of the passive margin succession on the northern (Timanian)
211 margin of the Baltic Shield in northern Fennoscandia. By integrating multiple lines of evidence
212 in the reconstruction, they establish that the entire succession is indigenous to Baltica and clarify
213 aspects of the margin's geometry up to the early stages of Timanian (late Neoproterozoic)
214 orogenesis. The DZ results also support a reliable baseline dataset documenting a Baltican
215 cratonic signature.

216

217 In their contribution, Anfinson et al. (2016) report on the temporal and spatial variations of
218 detrital zircon distributions in Permian and Triassic sandstones on the northern margin of
219 Canada's Sverdrup Basin. These data, together with existing sedimentological data, indicate that
220 sediments were derived from the recycling of remnant Ellesmerian orogenic deposits from a
221 source area of low relief to the north. However, the eastern segment of the basin margin received
222 a voluminous influx of Late Triassic sands containing abundant Early Carboniferous to Triassic
223 zircon populations, which they infer to be derived from the Taimyr region of Eurasia. This
224 localized influx is restricted to an interval of some 20 Myr. Integrating these results with
225 existing regional sedimentological and stratigraphic data, they further propose paleogeographic
226 reconstructions for the Permian and Late Triassic, the latter including a fluvial drainage system
227 connecting a Taimyr provenance with the Sverdrup Basin, circumventing the coeval Barents
228 Basin.

229

230 Ershova et al., (2016a), provide an overview of the Late Paleozoic paleogeography of Arctic
231 Russia based on a review of the Devonian to Permian facies belts of the Taimyr region, the major
232 island groups of the east Russian Arctic shelf and Chukotka. They recognize two distinct
233 paleogeographic groups. The widely separated island groups on the East Siberian shelf
234 (Severnaya Zemlya, New Siberian Islands, Wrangel) and Chukotka share stratigraphic and
235 provenance similarities but are distinctly different from the adjacent northern Siberian margin,
236 which suggests that Siberia experienced a distinctly different Late Paleozoic history. The
237 authors further suggest that the New Siberian Islands-Chukotka-Wrangel region may comprise
238 part of the distal Barents shelf, whereas the Kara terrane (Severnaya Zemlya-North Taimyr) was
239 accreted to Siberia in the Carboniferous.

240

241 A second contribution by Ershova et al. (2016b), documents an important baseline “Siberian
242 affinity” DZ signature, based on a study characterizing Siberia-derived Permian successions.
243 Permian strata deposited onto the Arctic margin of Siberia contain zircon age distributions that
244 are dominated by peaks of Late Carboniferous-Permian and Cambrian-Ordovician ages, with
245 minor Paleoproterozoic peaks indicating a cratonic source. On the basis of these baseline data,
246 the authors delineate two distinct lineages for Permian strata deposited into displaced terranes
247 within the Arctic realm. One lineage is dominated by zircon grains of Late Carboniferous to
248 Permian age similar to those studied from the northern Siberian margin and sourced from the
249 Ural-Mongolian orogen. The second lineage contains few or no grains of these ages and is
250 consistent with a source dominated by recycling of the Ellesmerian clastic wedge and Laurentian
251 margin rocks.

252

253 **3 Discussion**

254 *3.1. Large Igneous Provinces, Magmatic and Amagmatic Ocean Spreading*

255 Using the results gained from studies of areas with relatively well-understood tectonic settings,
256 such as the northern Atlantic and Eurasia basins, is it possible to gain new insights about areas
257 such as Amerasia Basin? Several papers in the Special Issue describe crustal structure and the
258 role of magmatic processes for areas near continent-ocean transitions along the east and north
259 Greenland margins, as well as within the Eurasian Basin. These results can be compared to new
260 data documenting the distribution of crustal types, including extensive areas of highly thinned
261 continental crust, as well as the effect of regional variations in the importance of magmatism in

262 the development of Canada Basin. Can these spatial variations perhaps be correlated with areas
263 of high-velocity magmatic underplating? Certainly, high-velocity underplated crust is identified
264 in the transitional zone surrounding the oceanic central part of Canada Basin, yet these areas
265 appear in areas identified as both magma-rich and magma-poor (Chain et al., 2016).

266

267 The Cenozoic tectonic history of the Arctic is dominated by the opening of the Eurasia Basin.
268 The initial separation of Lomonosov Ridge from the Barents Shelf and development of Gakkel
269 Ridge occurred ~55 Ma (e.g., Brozena et al., 2003; Schmidt-Aursch and Jokat, 2016). The early
270 history of mid-Atlantic Ridge propagation included extension and seafloor spreading in Labrador
271 Sea and Baffin Bay during separation of Greenland from North America, and subsequent
272 convergence with Ellesmere Island. A connection between the northward-propagating mid-
273 Atlantic Ridge and Gakkel Ridge was established only much later with the opening of Fram
274 Strait (Jokat et al., 2015; Hermann and Jokat, 2016).

275

276 In an ultra-slow spreading ocean-ridge system, mafic magmatism is apparently optional.
277 Amagmatic extension is developed in places along Gakkel Ridge (Schmidt-Aursch and Jokat,
278 2016). This study provides insights into the development and physical characteristics of slow-
279 spreading systems that may be instructive for other basins, in particular Amerasia Basin. Mafic
280 volcanism was locally important prior to Amerasia Basin initiation and throughout its subsequent
281 history, particularly in the northern part of the basin (compare Jokat et al., 2015, and Schmidt-
282 Aursch and Jokat, 2016 vs e.g., Evenchick et al., 2015). However, southern Canada Basin, is
283 identified as magma-poor (Chian et al., 2016).

284

285 The significance of large igneous provinces (LIPs) in the tectonic evolution of the Arctic has
286 been a topic of vigorous research over the past decade (e.g., Petrov et al., 2016). The
287 geodynamic drivers for development of LIPs and their tectonic consequences remain areas of
288 active research (e.g., Jokat et al., 2015; Hermann and Jokat, 2016; Anudu et al., 2016).

289

290 *3.2. Tectonics and Paleogeography*

291 Over the past decade, the Paleozoic paleogeographic histories of the Arctic region have become
292 a major focus of study. These activities have been manifest in three disparate fields, which are
293 nonetheless interconnected through various publication feedback cycles. These broad fields are:
294 1) tectonics and orogens; 2) continent-scale paleomagnetic syntheses; and 3) provenance and
295 paleogeography of displaced terranes (e.g., Cocks and Torsvik, 2007, 2011; Prokopiev et al.,
296 2013).

297

298 An improved understanding of the provenance and facies trends of displaced terranes and
299 deformed belts marginal to the major continents is critical to reconstructing paleogeographic
300 evolution. The relationships between Arctic terranes and the adjacent cratons are poorly
301 understood. However, new studies from the Arctic margins of Baltica, Russia and Canada
302 (Ershova et al., 2016a; Anfinson et al., 2016), as well as a comprehensive overview of the
303 Phanerozoic tectonic evolution of Alaska, comprising numerous displaced terranes inferred to
304 derive from both the Pacific and Arctic realms (Moore and Box, 2016), have provided much new

305 data and incorporate new results from multiple active research programs across the region. Their
306 new insights help to clarify the paleogeography of the Paleozoic and early Mesozoic strata across
307 the Arctic.

308

309 New detrital zircon U-Pb geochronology studies, some combined with trace element and rare
310 earth geochemistry, have provided significant insights into the provenance of Proterozoic to
311 Triassic strata (Zhang et al., 2016; Ershova et al., 2016b). These and other studies (e.g., Kirkland
312 et al., 2009; Anfinson et al., 2012; Lane and Gehrels, 2014) have provided baseline signatures for
313 autochthonous successions of known provenance, whereas others document exotic sources in
314 sedimentary successions (e.g., Rainbird et al., 1992; McNicoll et al., 1995; Miller et al. 2006;
315 Lane et al., 2016). Such studies suggest linkages between sedimentary source areas and remote
316 depositional basins have helped to inspire the theme of this issue. Improved understanding of
317 inter-regional linkages, whether sedimentary (provenance to basin), structural (correlation of
318 deformation events), or magmatic (in relation to both sedimentary and tectonic processes) will
319 improve Phanerozoic reconstructions of Arctic paleogeography.

320

321 *3.3. Progress and Problems*

322 Previously, alternative hypotheses of Arctic tectonic evolution were based on a paucity of data
323 and heavy reliance on speculation. Recently, due to a rapid increase in the volume of new
324 research, new geological models are based on more substantial datasets. Although the
325 geodynamic history of Eurasia Basin is well understood, several contributions to this volume
326 provide new results that refine our understanding of the interplay between magmatic ocean

327 spreading, non-magmatic spreading and the propagation of rifting. Segmentation of Gakkel ridge
328 into magma-rich and magma-poor domains is reflected in the crustal thickness and magnetic
329 signature of each domain. In the Amerasia Basin, similar issues may have been in play during its
330 formation in the Mesozoic. Delineation of the extent of oceanic crust is an important new
331 development, but so is the areal distinction of thinned continental lithosphere in some areas and
332 high-velocity magmatic underplated crust in other areas (Chian et al., 2016; cf. Li et al., 2016).
333 This makes clear that the role of magmatism extends beyond the question of age, extent and style
334 associated with Alpha-Mendeleev Ridge. There remains much to learn about the interplay of
335 magmatism and extension in the Amerasia Basin.

336

337 The Mesozoic tectonic evolution of the Amerasia Basin continues to be a topic of intense debate
338 with a variety of multi-stage hypotheses replacing earlier simple fan-shaped spreading scenarios
339 (e.g., Lane, 1997; Miller et al. 2006; Grantz et al. 2011; see also Lane et al., 2016, for
340 discussion). The documentation that oceanic crust is limited to the central part of the basin
341 (Chian et al., 2016) requires a reconsideration of the kinematics of basin formation and provides
342 further validation of the position of the extinct spreading axis within the oceanic part of the
343 basin.

344

345 Alpha-Mendeleev Ridge underlies much of the Amerasia Basin, yet it is understood largely
346 through remote sensing (seismic and potential field). While these data sources provide useful
347 structural and seismic velocity information, many of the results are ambiguous, leading to
348 differing, sometimes conflicting interpretations. Key among them is the question of its tectonic

349 setting and basement composition and origin (e.g., Petrov et al., 2016). Bruvoll et al. (2012)
350 provide a concise review of the literature concerning this problem. They conclude that “the
351 acoustic characteristics and seismic velocities compare more closely with basement on Ontong
352 Java Plateau and Kerguelen Plateau than normal ocean crust or wedges at volcanic margins”.
353 Oakey and Saltus (2016) provide new evidence in support of an interpretation that the Alpha-
354 Mendeleev basement consists of highly thinned and intruded continental crust, in agreement with
355 other recent studies (e.g., Lebedeva-Ivanova et al., 2006; Petrov et al., 2016).

356

357 Although Alpha-Mendeleev Ridge is regularly cited as a manifestation of the High Arctic Large
358 Igneous Province (HALIP), including in papers in this issue, the age range of Alpha-Mendeleev
359 Ridge is too poorly understood to be very confident of this correlation. The age of HALIP-
360 related magmatism in onshore exposures is widely cited as being concentrated in two main
361 pulses, one in Barremian-Aptian time and another in the Albian-Cenomanian (e.g., Evenchick et
362 al., 2015). Age constraints for Alpha-Mendeleev Ridge are rare. Jokat et al. (2013) report a
363 reliable Ar-Ar age of 89 ± 1 Ma on plagioclase from a tholeiite collected from central Alpha ridge.
364 This is more similar to ages from Hansen Point volcanics on Ellesmere Island (88-94 Ma; see
365 Anudu et al., 2016) and Peary Land dyke swarms on Greenland (80-85 Ma; see Jokat et al,
366 2015). The latter were cited as a possible early manifestation of Eurasia Basin initiation. Also,
367 variations in the chemistry of the magmas (alkaline vs tholeiite) have led some to suggest that
368 both plume and rifting environments may have been active (e.g., Jowitt et al., 2014). To the
369 south, in Labrador Sea and Davis Strait, long-lived magmatism includes significant episodes of
370 Late Jurassic to Early Cretaceous age well in advance of the dominant Paleogene events (Larsen
371 et al., 2009). Although much progress has been made, the questions of how many magmatic

372 events, what tectonic process(es) they were associated with, and what were their magma sources,
373 all remain largely unresolved. Suggestions that HALIP is a single entity spanning 130-70 Ma
374 (e.g., Petrov et al., 2016) may tend to cloud the issue rather than clarify it.

375

376 Recent technical advances in methodologies such as geochemistry and detrital zircon
377 geochronology have generated robust research programs to elucidate provenance and to
378 reconstruct paleogeography (Ershova et al., 2016b; Anfinson et al., 2016; and references
379 therein). However, the increasing recognition of zircon recycling indicates that the immediate
380 sources of distinctive age distributions may differ from the original magmatic sources (e.g., Lane
381 and Gehrels, 2011; 2014). Furthermore, although a distinctive zircon distribution may well be
382 derived from a particular source area, that does not imply that the source must be proximal to the
383 destination. In addition to recycling, the evidence for long distance transportation of zircon in
384 the sedimentary environment is well established (e.g., Rainbird et al., 1992; Anfinson et al.,
385 2016).

386

387 **4 Conclusions**

388 The elucidation of Arctic geodynamic and tectonic history has been a gradual process whereby
389 assumptions and speculations have been displaced by data and their interpretation. In this
390 Special Issue, interactions of lithospheric, tectonic, magmatic and sedimentation processes have
391 been documented and investigated. Each approach provides valuable information that can be
392 used to constrain elements of the region's tectonic history. However, a better understanding of
393 how these processes interact can provide additional insight that initially may not be apparent. By

394 presenting a suite of new research that includes studies of the mantle and whole crust we hope to
395 encourage broader consideration of multidisciplinary studies to help resolve refractory scientific
396 questions.

397

398 **5 Acknowledgements**

399 The Special Issue editors thank the contributors for their hard work and dedication in the
400 preparation of the papers presented here, and also Victoria Pease for her active support
401 throughout the process and in particular in co-convening the conference session giving rise to
402 this Special Issue. In particular, we thank the Editor-in-chief, Dr. Rob Govers for his patience,
403 guidance and valued advice throughout the process. Also, we appreciate the work of the
404 Tectonophysics editorial and production teams for bringing the Special Issue to print. R. Ernst,
405 G. Oakey and an anonymous reviewer provided a multitude of helpful suggestions to improve
406 the manuscript. This Special Issue is a contribution to the Geological Survey of Canada's
407 Geomapping for Energy and Minerals (GEM2) Program, Canada's Extended Continental Shelf
408 Program, and the Circum-Arctic Lithosphere Evolution (CALE) network. ESS Contribution
409 No. 20160152.

410 The Editors appreciate the important contributions of these Critical Reviewers:

411 A. Andreson
412 O. Anfinson
413 M. Cecile
414 V. Childers
415 S. Dehler
416 K. Dewing
417 G. Eagles
418 R. Ernst
419 G. Gehrels
420 P. Hart

421 D. Houseknecht
422 W. Jokat
423 A. Khudoley
424 F. Klingelhöfer
425 H. Lorenz
426 C. McFarlane
427 E. Miller
428 R. Mjelde
429 G. Oakey
430 T. Pavlis
431 K. Piepjohn
432 P. Potter
433 S. Roeske
434 G. Shellnutt
435 E. Solveig
436 C. Tegner
437 V. Verzhbitsky
438 K. Welford
439 and 18 Anonymous reviews.

440
441

442 **6 References**

443 Altenbernd, T., Jokat, W., Heyde, I. and Damm, V., 2016. Insights into the crustal structure of
444 the transition between Nares Strait and Baffin Bay. *Tectonophysics*, this issue.

445

446 Anfinson, O.A., Embry, A.F., and Stockli, D.F., 2016. Geochronologic and Petrochronologic
447 Constraints on the Permian-Triassic Northern Source Region of the Sverdrup Basin, Canadian
448 Arctic Islands. *Tectonophysics*, this issue.

449

450 Anfinson, O.A., Leier, A.L., Embry, A.F., and Dewing, K., 2012. Detrital zircon geochronology
451 and provenance of the Neoproterozoic to Late Devonian Franklinian Basin, Canadian Arctic
452 Islands. *Geol. Soc. Am. Bull.* 124, 415–430, doi: 10.1130/B30503.1

453

454 Anudu, G.K., Stephenson, R.A., Macdonald, D.I.M. and Oakey, G.N., 2016. Geological features
455 of the northeastern Canadian Arctic margin revealed from analysis of potential field data.
456 Tectonophysics, this issue.

457

458 Brozena J.M., Childers, V.A., Lawver, L.A., Gahagan, L.M., Forsberg, R., Faleide, J.I., and
459 Eldholm, O., 2003. New aerogeophysical study of the Eurasia Basin and Lomonosov Ridge:
460 Implications for basin development. *Geology* 31, 825-828.

461

462 Bruvoll, V., Kristoffersen, Y., Coakley, B.J., Hopper, J.R., Planke, S., Kandilarov, A., 2012. The
463 nature of the acoustic basement on Mendeleev and northwestern Alpha ridges, Arctic Ocean.
464 Tectonophysics, 514-517, 123-145. doi:10.1016/j.tecto.2011.10.015

465

466 Chian, D., Jackson, H.R., Hutchinson, D.R., Shimeld, J.W., Oakey G.N., Lebedeva-Ivanova, N.,
467 Li, Q. Saltus, R.W., and Mosher, D.C. 2016. Distribution of crustal types in Canada Basin,
468 Arctic Ocean. Tectonophysics, this issue.

469

470 Cocks, L.R.M., and Torsvik, T.H., 2007. Siberia, the wandering northern Terrane, and its
471 changing geography through the Paleozoic. *Earth-Sci. Rev.* 82, 29-74.
472 doi:10.1016/j.earscirev.2007.02.001

473

474 Cocks, L.R.M., and Torsvik, T.H., 2011. The Palaeozoic geography of Laurentia and western

475 Laurussia: A stable craton with mobile margins. *Earth-Sci. Rev.* 106, 1-51.

476 doi:10.1016/j.earscirev.2011.01.007

477

478 Ershova, V.B., Prokopiev, A.V., and Khudoley, A.K., 2016a. Devonian-Permian sedimentary

479 basins and paleogeography of the eastern Russian Arctic: An overview. *Tectonophysics*, this

480 issue.

481

482 Ershova, V.B., Khudoley, A.K., Prokopiev, A.V., Tuchkova, M.I., Fedorov, P.V., Kazakova,

483 G.G., Shishlov, S.B., and O'Sullivan, P.B., 2016b. Trans-Siberian Permian Rivers: A key to

484 understanding Arctic sedimentary provenance. *Tectonophysics*, this issue.

485

486 Evenchick, C.A., Davis, W.J., Bédard, J.H., Hayward, N., and Friedman, R.M., 2015. Evidence

487 for protracted High Arctic large igneous province magmatism in the central Sverdrup Basin from

488 stratigraphy, geochronology, and paleodepths of saucer-shaped sills. *Geol. Soc. Am.* 127, 1366-

489 1390. doi: 10.1130/B31190.1

490

491 Gac, S., Klitzke, P., Minakov, A., Faleide, J.I., and Scheck-Wenderoth, M., 2016. Lithospheric
492 Strength and elastic thickness of the Barents Sea and Kara Sea region. *Tectonophysics*, this
493 issue.

494

495 Grantz, A., Hart, P.E., and Childers, V.A., 2011, Geology and tectonic development of the
496 Amerasia and Canada Basins, Arctic Ocean, *in* Spencer, A.M., Embry, A.F., Gautier, D.L.,
497 Stoupakova, A.V., and Sorensen, K., eds., *Arctic Petroleum Geology*: Geological Society of
498 London Memoir 35, 501–508. doi: 10.1144 /M35.33

499

500 Hermann, T., and Jokat, W. 2016. Crustal structure off Kong Oscar Fjord, East Greenland:
501 Evidence for focused melt supply along the Jan Mayen Fracture Zone. *Tectonophysics*, this
502 issue.

503

504 Jakobsson, M., L. A. Mayer, B. Coakley, J. A. Dowdeswell, S. Forbes, B. Fridman, H.
505 Hodnesdal, R. Noormets, R. Pedersen, M. Rebesco, H.-W. Schenke, Y. Zarayskaya A, D.
506 Accettella, A. Armstrong, R. M. Anderson, P. Bienhoff, A. Camerlenghi, I. Church, M. Edwards,
507 J. V. Gardner, J. K. Hall, B. Hell, O. B. Hestvik, Y. Kristoffersen, C. Marcussen, R. Mohammad,
508 D. Mosher, S. V. Nghiem, M. T. Pedrosa, P. G. Travaglini, and P. Weatherall, 2012. The
509 International Bathymetric Chart of the Arctic Ocean (IBCAO) Version 3.0, *Geophysical*
510 *Research Letters*, doi: 10.1029/2012GL052219.

511

512 Jokat, W., 2003. Seismic investigations along the western sector of Alpha Ridge, Central Arctic
513 Ocean. *Geophys. J. Int.* 152, 185–201.

514

515 Jokat, W., Ickrath, M., and O'Connor, J., 2013. Seismic transect across the Lomonosov and
516 Mendeleev Ridges: Constraints on the geological evolution of the Amerasia Basin, Arctic Ocean.

517 *Geophys. Res. Lett.*, 40, 5047–5051, doi:10.1002/grl.50975.

518

519 Jokat, W., Lehmann, P., Damaske, D., and Nelson J.B., 2015. Magnetic signature of North-East
520 Greenland, the Morris Jesup Rise, the Yermak Plateau, the central Fram Strait: constraints for
521 rift/drift history between Greenland and Svalbard since the Eocene. *Tectonophysics*, this issue.

522

523 Jowitt, S.M., Williamson, M-C., and Ernst, R.E., 2014. Geochemistry of the 130 to 80 Ma
524 Canadian High Arctic Large Igneous Province (HALIP) Event and Implications for Ni-Cu-PGE
525 Prospectivity. *Econ. Geol.* 109, 281-307.

526

527 Kirkland, C.L., Pease, V., Whitehouse, M.J., and Ineson, J.R., 2009. Provenance record from
528 Mesoproterozoic-Cambrian sediments of Peary Land, North Greenland: Implications for the ice-
529 covered Greenland Shield and Laurentian palaeogeography. *Precambrian Research* 170 43–60,
530 doi: 10.1016/j.precamres.2008.11.006 .

531

532 Lane, L. S., 1997. Canada Basin, Arctic Ocean: Evidence against a rotational origin. *Tectonics*,
533 16, 363-387.

534

535 Lane, L.S., and Gehrels, G.S., 2011, Paleogeography and Provenance of Paleozoic rocks,
536 northwest Laurentia: A Story of Phanerozoic Zircon Recycling in northwest Canada. *In*: Stone,
537 D.B., Clough, J.G., and Thurston, D.H., eds., *Presentations made at International Conference*
538 *on Arctic Margins, ICAM 6, May 30–June 2, 2011, Fairbanks AK: Geophysical Institute Report*
539 *UAG-R-335, [http:// www2 .gi .alaska .edu /ICAMVI/](http://www2.gi.alaska.edu/ICAMVI/).*

540

541 Lane, L.S., and Gehrels, G.E., 2014. Detrital zircon lineages of late Neoproterozoic and
542 Cambrian strata, northwest Laurentia. *Geol. Soc. Am. Bull.* 126, 398–414,
543 doi:10.1130/B30848.1

544

545 Lane, L.S., Gehrels, G.E., and Layer, P.W., 2016. Provenance and paleogeography of the
546 Neruokpuk Formation, northwest Laurentia: An integrated synthesis. *Geol. Soc. Am. Bull.* 128
547 239-257. doi: 10.1130/B31234.1

548

549 Larsen, L.M., Heaman, L.M., Creaser, R.A., Duncan, R.A., Frei, R. And Hutchison, M., 2009.
550 Tectonomagmatic events during stretching and basin formation in the Labrador Sea and the

551 Davis Strait: evidence from age and composition of Mesozoic to Palaeogene dyke swarms in
552 West Greenland. *J. Geol. Soc. London* 166, 999-1012. doi: 10.1144/0016-76492009-038.

553

554 Lebedeva-Ivanova, N.N., Zamansky, Y.Ya., Langinen, A.E., and Sorokin, M. Yu., 2006. Seismic
555 profiling across the Mendeleev Ridge at 82°N: evidence of continental crust. *Geophys. J. Int.*
556 165, 527-544. Doi: 10.1111/j.1365-246X.2006.02859.x

557

558 Li, L., Stephenson, R.A., and Clift, P.D., 2016. The Canada Basin compared to the southwest
559 South China Sea: Two marginal ocean basins with hyper-extended continent-ocean transitions.
560 *Tectonophysics*, this issue.

561

562 McNicoll, V.J., Harrison, J.C, Trettin, H.P., and Thorsteinsson, R., 1995. Provenance of the
563 Devonian clastic wedge of Arctic Canada: Evidence provided by detrital zircon ages. *In:*
564 *Stratigraphic Evolution of Foreland Basins*, edited by S.L. Dorobek and G.M. Ross, SEPM
565 Special Publication 52, Society for Sedimentary Geology, Tulsa OK, USA, 77-93.

566

567 Miller, E.L., Toro, J., Gehrels, G., Amato, J.M., Prokopiev, A., Tuchkova, M.I., Akinin, V.V.,
568 Dumitru, T.A., Moore, T.E., and Cecile, M.P., 2006. New insights into Arctic paleogeography
569 and tectonics from U-Pb detrital zircon geochronology. *Tectonics* 25, TC3013, doi: 10

570 .1029 /2005TC001830.

571

572 Moore, T.E., and Box, S.E., 2016. Time Slice Maps Showing Age, Distribution and Style of
573 Deformation in Alaska North of 60°N: Implications for assembly of Alaska. *Tectonophysics*, this
574 issue.

575

576 Oakey, G.N., and Saltus, R.W., 2016. Geophysical analysis of the Alpha-Mendeleev Ridge
577 complex: Characterization of the High Arctic Large Igneous Province. *Tectonophysics*, this
578 issue.

579

580 Pease, V., Scarrow, J.H., Nobre Silva, I.G., and Cambeses, A., 2016. Devonian magmatism of
581 the Timan Range, Arctic Russia - subduction, post-orogenic, or rift related? *Tectonophysics*, this
582 issue.

583

584 Petrov, O., Morozov, A., Shokalsky, S., Kashubin, S., Artemieva, I.M., Sobolev, N., Petrov, E.,
585 Ernst, R.E., Sergeev, S., and Smelror, M., 2016. Crustal structure and tectonic model of the
586 Arctic region. *Earth-Sci. Rev.* 154, 29-71; doi.org/10.1016/j.earscirev.2015.11.013

587

588 Prokopiev, A.V., Ershova, V.B., Miller, E.L., and Khudoley, A.K., 2013. Early Carboniferous
589 paleogeography of the northern Verkhoyansk passive margin as derived from U–Pb dating of

590 detrital zircons: role of erosion products of the Central Asian and Taimyr–Severnaya Zemlya
591 fold belts. *Russian Geol. and Geoph.* 54, 1195-1204. <http://dx.doi.org/10.1016/j.rgg.2013.09.005>

592

593 Rainbird, R.H., Heaman, L.M., and Young, G.M., 1992. Sampling Laurentia: Detrital zircon
594 geochronology offers evidence for an extensive Neoproterozoic river system originating from the
595 Grenville orogen. *Geology*, v. 20, p. 351–354, doi: 10.1130 /0091-7613(1992)020

596

597 Schmidt-Aursch, M.C., and Jokat, W., 2016. 3D gravity modelling reveals off-axis crustal
598 thickness variations along the western Gakkel Ridge (Arctic Ocean). *Tectonophysics*, this issue.

599

600 United Nations, 1999. CLCS/11 - Scientific and Technical Guidelines of the Commission on the
601 Limits of the Continental Shelf; Commission on the Limits of the Continental Shelf (CLCS)
602 Selected documents of the Commission,

603 http://www.un.org/depts/los/clcs_new/commission_documents.htm (accessed 12 July 2016).

604

605 Zhang, W., Roberts, D. and Pease, V., 2015, Provenance of sandstones from Caledonian nappes
606 in Finnmark, Norway: Implications for Neoproterozoic-Cambrian palaeogeography.
607 *Tectonophysics*, this issue.

608

609 **Figure Caption**

610 Figure 1. Summary map showing the areal coverage of papers in this Special Issue. Outlines are
611 colour coded to correspond to the themes and topics outlined in the overview. Bathymetric base
612 map, IBCAO v3 is from Jakobsson, et al. (2012).

